

Copyright is owned by the Author of the thesis. Permission is given for a copy to be downloaded by an individual for the purpose of research and private study only. The thesis may not be reproduced elsewhere without the permission of the Author.

**Understanding Aspects of Andesitic Dome-forming Eruptions Through
the Last 1000 yrs of Volcanism at Mt. Taranaki, New Zealand**

A dissertation presented in partial fulfilment of the requirements
for the degree of
Doctor of Philosophy
in
Earth Science

Thomas Platz

Massey University, Palmerston North
New Zealand

2007



The summit of Mt. Taranaki. Photographed by G. Lube.

Dedicated to my parents,
to Katrin and my son August

Abstract

Andesitic volcanoes are notorious for their rapid and unpredictable changes in eruptive style between and during volcanic events, a feature normally attributed to shallow-crustal and intra-edifice magmatic processes. Using the example of eruptions during the last 1000 yrs at Mt. Taranaki (the Maero Eruptive Period), deposit sequences were studied to (1) understand lava dome formation and destruction, (2) interpret the causes of rapid shifts from extrusive to explosive eruption styles, and (3) to build a model of crustal magmatic processes that impact on eruption style.

A new detailed reconstruction of this period identifies at least 10 eruptive episodes characterised by extrusive, lava dome- and lava flow-producing events and one sub-Plinian eruption. To achieve this, a new evaluation procedure was developed to purge glass datasets of contaminated mineral-glass analyses by using compositional diagrams of mineral incompatible-compatible elements. Along with careful examination of particle textures, this procedure can be broadly applied to build a higher degree of resolution in any tephrostratigraphic record. Geochemical contrasts show that the products of the latest Mt. Taranaki eruption, the remnant summit dome (Pyramid Dome) was not formed during the Tahurangi eruptive episode but extruded post-AD1755. Its inferred original maximum volume of $4.9 \times 10^6 \text{ m}^3$ (DRE) was formed by simultaneous endogenous and exogenous dome growth within days. Magma ascent and extrusion rates are estimated at $\geq 0.012 \text{ ms}^{-1}$ and $\geq 6 \text{ m}^3 \text{ s}^{-1}$, respectively, based on hornblende textures. Some of the Maero-Period dome effusions were preceded by a vent-clearing phase producing layers of scattered lithic lapilli around the edifice [Newall Ash (a), Mangahume Lapilli, Pyramid Lapilli]. The type of dome failure controlled successive eruptive phases in most instances. The destruction of a pressurised dome either caused instantaneous but short-lived magmatic fragmentation (Newall and Puniho episodes), or triggered a directed blast-explosion (Newall episode), or initiated sustained magmatic fragmentation (Burrell Episode). The transition from dome effusion to a sustained, sub-Plinian eruption during the Burrell Lapilli (AD1655) episode was caused by unroofing a conduit of stalled magma, vertically segregated into three layers with different degrees of vesiculation and crystallisation. The resultant ejecta range from brown, grey and black coloured vesicular clasts to dense grey lithics. Bulk compositional variation of erupted clasts can be modelled by fractionation of hornblende, plagioclase, clinopyroxene, and Fe-Ti oxides. Pre-eruption magma ascent for the Maero Period events is assumed to begin at depths of c.9.5 km.

Acknowledgements

I wish to thank every person who contributed to the outcome of this study.

I went back to New Zealand to study one of the most fascinating phenomena nature has to offer: volcanoes. Of the North Island volcanoes, Mount Taranaki stands out as being the most imposing, and to me the most striking. Despite its dormancy, it proved a real challenge, to climb, find samples and sections, and to draw out some of his secrets.

For the opportunity to work again on Mt. Taranaki I am indebted to my chief-supervisor Dr Shane J. Cronin. Through him I learnt much about how to observe volcanic deposits and to understand the various processes involved in generating them. His aid in my receiving a Massey University Doctoral scholarship is highly appreciated. In past years Dr Cronin supported numerous overseas trips, which offered me the opportunity to meet other scientists, either on conferences or at the institutions where my overseas supervisors are based. I also benefited from his extraordinary skills in writing and presenting ideas and thoughts.

I also wish to thank all my co-supervisors: Prof. Vince E. Neall and Dr Robert B. Stewart (Massey University, New Zealand), Prof. Katharine V. Cashman (University of Oregon, USA), and Prof. Stephen F. Foley (University of Mainz, Germany) for their guidance and support over the years, their fruitful discussions and criticisms. I gratefully acknowledge the skill of all my supervisors to catch my thoughts especially when articulated in a Germanic English.

During my studies a lot of people were involved that helped and supported me from different parts of the globe.

New Zealand: First of all I wish to thank Jonathan N. Procter, Michael B. Turner, Anke V. Zernack and Kat A. Holt, my fellow inhabitants in the “Magma Chamber”. Without their friendship, life at Massey would have been colourless. Sharing “ups and downs” made many situations easier to bear.

I am also grateful to the staff of the Soil and Earth Sciences Group for assisting with everyday needs: Bob Toes (for also helping to fix Emma), Ian F. Furkert, Ross W. Wallace, Mike R. Bretherton, Moira Hubbard, Glenys C. Wallace, Anne R. West, Dr R. Clel Wallace, Dr Jerome A. Lecointre, Julie A. Palmer and Dr Alan S. Palmer. Carolyn B. Hedley (Landcare Research) introduced me to the thermal analysis technique. Dr Ritchie Sims and John Wilmshurst assisted in XRF- and electron microprobe analysis at The University of Auckland. For discussions on aspects of Taranaki and access to their

geochemical data I sincerely thank Assoc. Prof. Ian E.M. Smith (The University of Auckland) and Prof. Richard C. Price (Waikato University). Garry Bastin (Taranaki Research Centre) is thanked for his assistance in looking for historic references.

Eugene: From the Department of Geological Sciences I wish to thank Dr Heather M.N. Wright for introducing and assisting in sample preparation and measurement using the He-pycnometer and porosimeter. For inspiring discussions on pumice I especially acknowledge Drs Alison C. Rust and Heather M.N. Wright. Emily Johnson, Dr Julie Roberge, and Celeste Mercer are also thanked for their help in preparing melt inclusions. For discussions and access to the FTIR- and SEM facilities, the help of Assoc. Prof. Paul J. Wallace and Dr John J. Donovan are appreciated.

Kiel: Encouraging and fruitful discussions with Dr Armin Freundt and Sebastian Münn (IfM-GEOMAR) are especially acknowledged.

Mainz: For assistance in sample preparation and measurement I wish to thank the team on the Laser-ICP-MS, Institute of Geosciences, especially Drs Franziska Nehring and Matthias Barth.

Copenhagen: For encouragement, discussion, and analysis Dr Holger B. Lindgren of the Geological Survey of Denmark and Greenland is highly appreciated.

Potsdam: My thanks go also to Dr Thomas R. Walter (GeoForschungsZentrum Potsdam) for his unreserved discussions.

Studying abroad would have been far more difficult without the support of friends and family back home. My sincere thanks go to my parents who always supported me in everything I pursued. My friends Dr Gert Lube, Sebastian Münn, Dr Franziska Nehring, Luzie Herklotz, Dr Diana Reckien, and Jonathan N. Procter were always there when I needed them.

But the most important part was contributed by Katrin Bauer, through her love, respect, and honesty.

This work was funded and supported by a Massey University Doctoral Scholarship, the George Mason Trust of Taranaki, the Helen E. Akers Scholarship, the FRST-PGST contract MAUX0401, and an Institute of Natural Resources transitional scholarship.

Table of Contents

Abstract	i
Acknowledgements	iii
List of Tables	xi
List of Figures	xiii
Chapter 1 Introduction	1
1.1 Introduction	1
1.2 Objectives and Strategy	5
1.3 Thesis Outline	6
1.4 Background Geology	7
1.4.1 Regional Geological Setting	7
1.4.2 Taranaki Basin	9
1.4.3 Taranaki Volcanic Lineament	12
1.4.4 Mount Taranaki/Egmont	12
1.4.5 Petrology of Mt. Taranaki Rocks	15
1.5 References	19
Chapter 2 Methodology	31
2.1 Brief Outline	31
2.2 Field Studies	32
2.3 Mineralogy	33
2.3.1 Sample Preparation	33
2.3.2 Microscopic Studies	34

2.4	Geochemistry	34
2.4.1	X-ray Fluorescence Spectroscopy	34
2.4.2	Electron Microprobe Analysis	35
2.4.3	Laser Inductively Coupled Plasma Mass Spectrometry	36
2.5	Porosity and Permeability	37
2.6	Scanning Electron Microscopy	41
2.7	Fourier Transform Infrared Spectroscopy	41
2.8	Thermal Analysis	43
2.9	References	46
Chapter 3	The Maero Eruptive Period	49
3.1	Introduction	49
3.1.1	Previous Studies	51
3.1.1.1	Previous Work on Lava Flow Stratigraphy	54
3.2	Results	57
3.2.1	Stratigraphic Type and Reference Sections	57
3.2.1.1	Type Section of the Maero Formation	58
3.2.1.2	Reference Sections of the Maero Formation	64
3.2.2	Block-and-Ash Flow Deposits	68
3.2.2.1	Distribution and Flow Paths	68
3.2.3	Lava Flows	71
3.2.4	Glass Chemistry	75
3.2.4.1	Special Characteristics of Some Erupted Units	76
3.2.4.2	Correlation of Tephra and Pyroclastic Flow Deposits	77
3.3	Discussion	81
3.3.1	Eruption Frequency of the Maero Eruptive Period	81
3.3.2	Comparison to the Previously Known Stratigraphy	86
3.3.3	Tephrostratigraphy of the Maero Eruptive Period	87
3.4	References	91

Chapter 4	Improving the Reliability of Microprobe-based Glass Analyses	95
4.1	Abstract	97
4.2	Introduction	97
4.2.1	Andesitic Volcanism, Tephra Generation and Dispersal	100
4.2.2	Sample Sites	103
4.3	Results	103
4.3.1	Contrasts in Particle Texture	103
4.3.2	Glass Chemistry and Data Evaluation	105
4.3.3	Estimating Plagioclase Proportions in Contaminated Analyses	111
4.4	Discussion	112
4.5	Conclusions	116
4.6	References	117
Chapter 5	Non-explosive, Dome-forming Eruptions	121
5.1	Introduction	121
5.1.1	Lava Domes	122
5.1.1.1	Lava Dome Distribution in Taranaki	124
5.2	Field Observations	124
5.2.1	Tahurangi Eruptive Deposits	124
5.2.2	The Present Summit Lava Dome	128
5.2.3	Rock-Avalanche Deposit	130
5.3	Sample Sites and Methods	131
5.4	Results	132
5.4.1	Dome Volume Calculations	132
5.4.2	Mineralogy and Mineral Chemistry	134
5.4.3	Bulk Rock Composition	139
5.4.4	Microstructure, Density and Permeability of Dome Rocks	142
5.5	Discussion	142

5.5.1	Lava Dome Emplacement and Growth	142
5.5.2	Lava Dome Collapse	145
5.5.3	Estimation of Eruption Parameters	148
5.5.3.1	Magma Source Areas	148
5.5.3.2	Magma Storage at the Level of Neutral Buoyancy	151
5.5.3.3	Inferences for Hornblende Stability	153
5.5.3.4	Magma Ascent Rate	154
5.5.4	Eruption Duration	155
5.5.5	Approximation of the Time of Eruption	158
5.5.6	Historic References to Volcanic Activity in the 19 th Century	161
5.6	Conclusions	166
5.7	References	167
Chapter 6	Transition to Explosive Eruptions	175
6.1	Abstract	177
6.2	Introduction	178
6.2.1	Geological Setting	179
6.3	Methodology	180
6.4	Results	182
6.4.1	Field Observations	182
6.4.1.1	Fall Deposits	182
6.4.1.2	Pumice Pyroclastic Flow Deposits	182
6.4.1.3	Block-and-Ash Flow Deposits	185
6.4.2	Volume Estimates	185
6.4.3	Clast Types and Textures	186
6.4.4	Mineralogy and Mineral Chemistry	189
6.4.5	Bulk Rock Geochemistry	190
6.4.6	Glass Composition	194
6.4.7	Porosity and Permeability	196

6.5	Discussion	198
6.5.1	Pre-climactic Conditions	199
6.5.1.1	Grey Lithics	199
6.5.1.2	Grey Pumice (Unit 1)	201
6.5.1.3	Brown Pumice (Unit 2)	202
6.5.2	Syn-climactic Conditions	203
6.5.2.1	Black, Banded, and Unit 3 Grey Pumice Clasts	203
6.5.3	Eruption Dynamics	204
6.6	Conclusions	206
6.7	References	208
6.8	Appendix	213
6.8.1	Appendix A: Correlation of Burrell Deposits, NW Sector	213
6.8.2	Appendix B: Location of Pyroclastic Flow Deposits	214
6.8.3	Appendix C: Volatile Contents in Glass – Preliminary Results and Conclusions	215
6.8.3.1	FTIR Spectroscopy	215
6.8.3.2	Thermal Analysis	217
Chapter 7	Reconstruction of Eruption Mechanisms Using Physico-chemical Data	221
7.1	Introduction	221
7.1.1	Viscosity of Magmas	222
7.2	Methods and Approach	225
7.3	Results	225
7.3.1	Petrography	225
7.3.1.1	Pyroclastic Flow Deposits	225
7.3.1.2	Lava Flows	230
7.3.2	Bulk Rock Chemistry	230
7.3.2.1	Pyroclastic Flow Deposits	230
7.3.2.2	Lava Flows	232

7.3.3	Physical Properties	232
7.3.3.1	Water Estimates of Volcanic Glasses	233
7.3.3.2	Viscosity	234
7.4	Discussion	238
7.4.1	Comparison of Physico-chemical Properties	238
7.4.1.1	Bulk Rock Composition	238
7.4.1.2	Melt and Magma Viscosities	239
7.4.2	Course and Eruption Styles of the Maero Eruptive Period	243
7.4.2.1	Types of Lava Domes	243
7.4.2.2	Causes of Dome Collapse and Associated Deposits	244
7.4.2.3	Reconstruction of the Maero Eruptive Period	247
7.5	Conclusion	251
7.6	References	252
Chapter 8	Conclusions	259
8.1	Avenues of Future Research	263
Appendices		I

List of Tables

Table 2-1.....	36
Detection limit of element oxides measured at the EMP (University of Auckland) including the deviation from a reference glass composition.	
Table 2-2.....	43
Water and carbon dioxide peaks on the FTIR spectra including their bonds within glasses.	
Table 3-1.....	52
Stratigraphy of the youngest deposits of Mt. Taranaki (Druce, 1966).	
Table 3-2.....	53
Stratigraphy of the youngest deposits at Mt. Taranaki until 2003 (Druce, 1966; Neall, 1972; Neall, 1979; McGlone et al., 1988; Lees and Neall, 1993).	
Table 3-3.....	54
Stratigraphy of the last 1000 yrs of activity at Mt. Taranaki established in 2003 (modified after Cronin et al., 2003).	
Table 3-4.....	73
Sample list, description and stratigraphy of studied lava flows and scoria-and-ash flow deposits.	
Table 3-5.....	83
New tephrostratigraphy of the Maero Eruptive Period.	
Table 4-1.....	106
Glass EMPA of Taranaki (Burrell Lapilli eruption) and Ruapehu (14. October 1995 eruption) sorted by classed uncontaminated and contaminated data points. Detection limits for Taranaki glass EMPA only. All data in wt. %.	
Table 4-2.....	108
Plagioclase microlite rim and centre compositions (Burrell Lapilli eruption, Mt. Taranaki). All data in wt. %.	
Table 4-3.....	112
Calculation of various plagioclase mixing proportions within Taranaki hybrid glass EMPA. EMPA data in wt. %. See text for explanation.	
Table 5-1.....	132
Sample location and description of studied summit dome rocks.	
Table 5-2.....	134
Comparison of the modelled summit dome to other Taranaki flank lava domes.	
Table 5-3.....	156
Historic lava dome eruptions compiled from Newhall and Melson (1983) and the Smithsonian Institution Catalogue.	
Table 5-4.....	158
Physical parameters and results for equations 5-4 to 5-6.	
Table 5-5.....	160
Physical parameters used for conductive cooling and rainfall-quenching of the lava dome.	

Table 5-6.....	163
Selected eyewitness reports of the 18 th and 19 th century.	
Table 6-1.....	186
Minimum volume and mass calculations of erupted tephra.	
Table 6-2.....	191
XRF bulk rock geochemistry of Burrell Lapilli clasts. Samples are mostly from pumice flow deposits with one BAF deposit sample (P10).	
Table 6-3.....	193
Mass balance calculation for the inferred fractionation process for pumice-lithics.	
Table 6-4.....	195
Representative normalised matrix glass compositions. All data in wt. %.	
Table 6-5.....	214
Location of pyroclastic flow deposits on the NW and S sector of Mt. Taranaki.	
Table 7-1.....	226
Sample list and description of studied pyroclastic flow deposits.	
Table 7-2.....	228
Summary of petrographic studies of BAF samples.	

List of Figures

Figure 1-1.....	8
Tectonic setting and its relation to Holocene volcanism in the North Island, New Zealand. Area between the Hikurangi Trough and the axial ranges is the forearc. Arrow indicates Pacific Plate movement with a rate of 42 mm yr ⁻¹ . AVF – Auckland Volcanic Field; TgVC – Tongariro Volcanic Centre; TVL – Taranaki Volcanic Lineament; TVZ – Taupo Volcanic Zone. AD and RD refers to andesite and rhyolite dominance within TVZ, respectively. Modified after Reyners et al. (2006) and Wilson et al. (1995).	
Figure 1-2.....	10
Regional tectonic setting of the Taranaki peninsula. The Taranaki Volcanic Lineament comprises the Sugar Loaf Islands (SLI), Kaitake (K), Pouakai (P) and Mt. Taranaki (T). Major onshore (thick lines) and offshore (thin lines) faults are indicated: IF-Inglewood Fault, MF-Manaia Fault, NF-Norfolk Fault, OF-Oaonui Fault. Contours are at 300 m intervals for the volcanic edifices only. New Plymouth (NP) as the major settlement is identified. Modified after Sherburn and White (2005).	
Figure 1-3.....	14
Orthophotograph of Mt. Taranaki including the lower northwestern flanks. Inset shows upper cone area including Fanthams Peak. Major morphological features are indicated by letters: a-summit crater of Mt. Taranaki, b-Fanthams Peak, c-the Beehives (two lava domes), d-scarp of the Opuia amphitheatre, e-Big Pyramid, f-The Dome, g-Skinner Hill (probably a buried dome structure), h-Pyramid Stream, i-Maero Stream, j-Waiweranui Stream, k-Hangatahua River, l-Egmont National Park boundary (forest/pasture border, here 390 m), m-Kapuni Gorge (marks the eastern border of the amphitheatre), n-Sharks Tooth (second highest peak, 2510 m), o-Fanthams Peak (comprising multiple vents), p-remnant summit dome, q-Turtle, r-Bobs Ridge (western border of the Opuia amphitheatre), s-NW flank and main path for BAFs.	
Figure 1-4.....	16
Classification of Mt. Taranaki and Mt. Ruapehu rocks after Gill (1981). Modified after Price et al. (1999).	
Figure 1-5.....	17
Trace element patterns of selected Mt. Taranaki rocks. Warwicks/Staircase refers to lava flow groups of the main cone of Mt. Taranaki; Fanthams describes lava flows of the satellite vent Fanthams Peak. Data compiled from Price et al. (1992, 1999). Normalisation after Sun and McDonough (1989).	
Figure 2-1.....	38
The specimen illustrates how pumice cores were obtained. As in this case, six cores in three mutually perpendicular orientations were drilled.	
Figure 2-2.....	39
Standard deviations for diameter and length of the cores (a) and for volumes V_c and V_{He} (b). It is noted that one sample in a) is off the chart at a standard deviation of 0.1321 cm. b) Samples are differentiated into those drilled in Oregon and at Massey. Oregon samples show small variations for both V_c and V_{He} .	
Figure 2-3.....	40
Standard deviation for connected porosities. The limit of 0.3 cm ³ for V_c and V_{He} in Fig. 2-2b is used as maximum limit. Cores drilled in Oregon are shown for comparison.	
Figure 2-4.....	40
Permeability measurements of individual cores were performed three to six times, partially using multiple flow rates, in order to assess reproducibility. In this case, the red graph suggests higher flow rates compared to the other three runs and was excluded from further calculations.	
Figure 2-5.....	44
Separated groundmass glass fraction from a pumice clast (SD32).	

Figure 2-6.....	45
Thermal analysis of volcanic glass to 850 °C. The change in weight (green axis) was measured after the isothermal break at 110 °C.	
Figure 3-1.....	55
Distribution of lava flows on the Mt. Taranaki main cone and Fanthams Peak. Sampled lava flows and scoria-and-ash flow units are highlighted. The thick solid line represents the Opuā amphitheatre scarps. The map has been modified after Neall (2003); map details were kindly provided by J.N. Procter.	
Figure 3-2.....	57
Location of type and reference sections around the edifice of Mt. Taranaki. TS-type section, P-Pembroke Road, W-Waingongoro Stream, M-Manaia Road. Numbered black squares refer to reference sections. Black diamonds mark distal locations of pyroclastic flow deposits in the area of Saunders Road – Wiremu Road – Waiweranui Stream.	
Figure 3-3.....	59
Type section of the Maero Formation located in Maero Stream, NW sector of Mt. Taranaki. Reference sections Nos. 1 and 2 are located in Pyramid Stream and Hangatahua River, respectively. The outcrops comprise deposits of BAFs, surges and pumice flows as well as co-ignimbrite ashes. Debris flow-, lahar-, and fluvial deposits are also exposed but not always differentiated. For further details see text and appendices. Labels Py7 etc refer to sample numbers. See also Fig. 3-4 for field photographs.	
Figure 3-4.....	62
Field photographs of deposits from the NW sector of Mt. Taranaki. a) NW sector of Mt. Taranaki as viewed from the summit. Note the depression in vegetation caused by an avalanche (see Chapter 5). H. R. - Hangatahua River . b) exposure in the upper right Pyramid fork. Cliff section is approx. 30 m high and shows mainly lahar and hyperconcentrated flow deposits. Note Shane Cronin for scale (arrow). c) outcrop on the true left side of the Maero Stream near the intersection of Puniho and Holly Hut track. The upper unit shows the Taurangi Block-and-Ash Flow deposit (unit a) with its upper matrix-supported and lower clast-supported zones. d) reference section No.1, lower Pyramid Stream, true right side. Only major units are labelled. Discolouration of units IV-Bb and II-B is caused by a raised iron-rich water table. Note the black bar has same height as the spatula (arrow). e) close-up of c); basal portion of Taurangi BAF (unit a) which consists of fine to medium ash. Dashed line marks the boundary between main body of the BAF and its basal portion. f) close-up of d); finely laminated and cross-bedded fine to medium ash surge deposit. Some pumice clasts (arrows) are present. g) exposure on the true right side of Maero Stream. Major units are labelled. The reworked section is 2.5 m thick and represents predominantly fluvial deposits. h) close-up of g); unit IV-Bb. Noteworthy are boulder trains and weak reverse grading of the lower to middle portion. i) distal blast deposit [Newall Breccia (a)] showing its typical pocketing appearance. j) stream exposure of a BAF deposit with pervasive red coloured top portion and grey bottom portion (Waiweranui Stream). k) in situ charcolised tree at a distal BAF exposure [Burrell Breccia (A)] near Saunders Road. l) degassing pipe originating at a charred log within the BAF deposit (same unit as in k). Note that the pipe branches at the centre of the photograph. Same unit as in k). Labelled units in c, d and g refer to the lithostratigraphic code. See discussion of Chapter 3 for further details. Detailed description of reference section No.1 can be found in the appendix. Photographs of f) and i) were taken by Shane Cronin.	
Figure 3-5.....	64
Reference sections Nos. 3 and 4 located on the east flank of Mt. Taranaki. Reference site: No.3–Pembroke Road cutting; No.4–intersection of Wain- gongoro Stream and Round-The-Mountain-Track. E02-72 etc are sample numbers.	

Figure 3-6.....	66
Reference sections Nos. 5 and 6 located on the south flank of Mt. Taranaki. Reference site: No.5--at the site of the Old Mangahume Hut; No.6--near the Old Mangahume Hut, upslope of No.5. Labels E03-35 etc are sample numbers.	
Figure 3-7.....	67
Reference sections Nos. 7, 8 and 9 located on the west and north sector of Mt. Taranaki. Reference site: No.7--Ahukawakawa Swamp; No.8--Marupakoko Stream near Kahui Hut; No.9--Parihaka Road cutting. Labels E02-4 etc refer to sample numbers.	
Figure 3-8.....	68
Crater rim stratigraphy as exposed on its SW side at the entrance of Okahu Gorge. Four lava flows are identifiable with the youngest flow (4) being known as South Flow.	
Figure 3-9.....	69
Composite photograph of the upper NW flank of Mt. Taranaki showing named and described morphological features.	
Figure 3-10.....	70
Flow paths and distribution of BAF and surge deposits on the N to W sectors. Crater exit areas: 1-north sector, 2-northwest sector, 3-southwest sector.	
Figure 3-11.....	75
Histogram of all glass chemical analyses for SiO ₂ (a) and K ₂ O (b).	
Figure 3-12.....	77
Backscatter-electron microscopy images of individual glass shards. Different glass shard shapes within individual samples are illustrated (top: vesicular, partially deformed; bottom: dense and angular). Label in images gives analysis number and approximate beam location (black dot). Grey=glass; light grey=minerals; black=epoxy. a-b) Tahurangi Ash sample (T04-98) with chemically homogenous glass shards but of different texture; a) deformed vesicles, b) dense angular. Note that the presence of large minerals (bottom) may alter vesicle distribution. c-d) Newall Ash sample (E02-75) with distinct shard textures; glass chemistry of shard (d) is similar to those of scoria-and-ash flow glass shard (c). Note regular to irregular vesicles in c). e-f) Burrell Lapilli sample (E03-35) with Taranaki glass shard (f) and Taupo volcano-derived glass shard (e), which shows a deformed vesicle.	
Figure 3-13.....	78
Glass chemistry of Burrell Lapilli-erupted deposits. The SiO ₂ vs. K ₂ O diagram shows a general positive correlation. Highest mean silica and potassium contents are observed for BAF and surge deposits. The crosses represent the 95% confidence interval of the sample mean. The tephra sample E03-35 is bimodal containing glass shards with a signature similar to Taupo volcano (or atypical of Taranaki). If foreign glass shard analyses are excluded the sample mean is located within the field of BAF/surges (=E03-35'). Also included are sample means of a surge deposit pre-dating the Maero Eruptive Period showing a rhyolitic glass chemistry. It is noted that one sample (WW19) shows large variations and a bimodal sample population.	
Figure 3-14.....	79
Comparison of mean sample values of scoria-and-ash flow units (T04-53, T04-56) to pyroclastic pumice flow deposits of the Burrell episode (units 1-3) and Puniho Ash, and other Maero deposits (crosses). Individual glass shards (small black squares) within deposits, other than T04-53 and T04-56, that have compositions similar to scoria-and-ash flows. a) SiO ₂ vs. K ₂ O, b) CaO vs. FeO.	

Figure 3-15.....	80
Correlation of individual tephra and pyroclastic flow units based on field studies and glass chemistry (SiO ₂ vs. K ₂ O). a) Waingongoro and Waiweranui episodes, b) Newall and Puniho episodes, c) Taurangi and Te Popo episodes.	
Figure 4-1.....	99
Locations of Holocene volcanic centres in the North Island of New Zealand: Auckland Volcanic Field, Taupo Volcanic Zone (TVZ), Tongariro Volcanic Centre (TgVC), Mt. Taranaki. Numbers 1-4 refer to distal andesite tephra sites: 1-Onepoto Basin/Pukaki Lagoon/Lake Pupuke (Sandiford et al., 2001; Shane and Hoverd, 2002; Shane, 2005), 2-Waikato lakes (Lowe, 1988b), 3-Kaipu Bog (Lowe et al., 1999), 4-Lake Tutira (Eden and Froggatt, 1996), 5-Lake Poukawa (Shane et al., 2002), 6-Kaimanawa Mts. and Ruahine Ranges (Froggatt and Rogers, 1990).	
Figure 4-2.....	104
Particle textures found in Burrell Lapilli sub-Plinian fall deposits. a) pumice clast, type 1, with clear to pale brownish glass; b) hypocrySTALLINE groundmass texture of type 2 clast showing plagioclase, Fe-Ti oxide and minor clinopyroxene microlites; c) semi-vesicular type 3 clast with brown groundmass glass, large crystals are hornblende; d) for comparison, hypocrySTALLINE groundmass of the present summit dome of Mt. Taranaki with abundant microlites of plagioclase, Fe-Ti oxides and minor clinopyroxene. Scale bars are in µm.	
Figure 4-3.....	105
Standard deviations of all major oxides are shown for the original glass EMPA dataset and the glass dataset classed as uncontaminated by plagioclase for the Taranaki (a) and Ruapehu (b) samples. A clear reduction in SiO ₂ , Al ₂ O ₃ and CaO variations is observed. See text for further details. All Fe expressed as FeO.	
Figure 4-4.....	109
Bivariate plots of Al ₂ O ₃ and K ₂ O vs. SiO ₂ for Burrell Lapilli (Taranaki) and Ruapehu glass data. Data points with lowest SiO ₂ show linear relationship towards mean plagioclase compositions (dashed lines).	
Figure 4-5.....	110
Bivariate oxide plots for Burrell Lapilli eruption, Taranaki (a-d) and 14. October 1995 eruption, Ruapehu (e-f) demonstrate how contaminated glass analyses were identified. Open symbols are classed uncontaminated, closed symbols represent hybrid glass-plagioclase analyses. Dashed lines point towards mean plagioclase compositions.	
Figure 4-6.....	114
The glass evaluation procedure cannot be directly applied to BAF deposits as demonstrated for the Burrell Lapilli equivalent BAF deposit. Although data points above a threshold value of 17.9 wt.% Al ₂ O ₃ clearly embrace contaminated glass analyses, the transitional data between 17.1 and 17.9 wt.% Al ₂ O ₃ cannot be uniquely classified. Dashed line points toward mean plagioclase composition.	
Figure 4-7.....	115
Glass compositions of Taranaki and TgVC tephtras (Lowe 1988b; Lowe et al., 1999; Eden and Froggatt, 1996 and Shane and Hoverd, 2002) show large variations, here only shown the means and standard deviations for K ₂ O and Al ₂ O ₃ . Contaminated plagioclase-glass analyses and/or the analysis of two or more particle types may have caused the apparent glass compositional heterogeneity. The small variation in the unmodified Burrell Lapilli dataset (Taranaki) is shown for comparison (in grey).	
Figure 5-1.....	123
Lava dome types: a) spiny (Rock Mesa ENE, Oregon), b) spiny-lobate (Mt. St. Helens lava dome, Washington, July 2004), c) lobate-platy (Big Obsidian Flow, Newberry, Oregon).	

Figure 5-2.....	125
NW sector of Mt. Taranaki showing the main deposition area of Maero BAF deposits. The extent of the Tahurangi BAF A is outlined (light grey); unit B is only observed in the Maero Stream and is omitted for clarity. The rock-avalanche deposit (RAD) is outlined as observed on aerial photographs from 1959 (mid-grey) with an additional area based on field observations (dark grey). Contour interval is 20 m. The upper right inset shows the general slope inclinations.	
Figure 5-3.....	127
Correlation of Tahurangi BAF A and B units and the rock-avalanche deposit across the Pyramid-Maero-Hangatahua area. The exposures are sorted by stream and planimetric distance from source (filled squares in Fig. 5-2). See figure legend for further details (RAD – rock-avalanche deposit). Magnetism was measured by a portable fluxgate magnetometer. For clarity, older exposed units were omitted. Outcrop numbers and profiles refer to Platz (2001) except S04-133.	
Figure 5-4.....	129
The remnant present summit dome. a) hemispherical shape of the dome as viewed from the SE; arrow points to a person for scale. Photographed by S.J. Cronin. b) dome amphitheatre as viewed from the W; the arrow points to the summit marker (2518 m); the dashed lines mark the hydrothermally altered central dome portion. c) northern scar of the amphitheatre showing listric faults. d) the summit marker is a sub-vertical extrusion penetrating the carapace; note weak columnar jointing; summit marker is highest point of the dome. e) orthophotograph of the summit region of Mt. Taranaki; outline shows mapped deposits associated with the lava dome; note the blocky lava flow to the N; crosses mark sample locations. f) the ‘Three Sisters’ (background) mark the NW border of the intra-crater collapse zone with resulting deposit still preserved in the crater (foreground).	
Figure 5-5.....	130
Black and white photograph taken between 1898 and 1901 showing the fresh bouldery rock-avalanche deposit (centre to right). The photograph is taken from the Round-The-Mountain-Track just west of Maero Stream from the top of a buried lava ridge. The view is NNW towards Pouakai. Photographed by the surveyor H.M. Skeet.	
Figure 5-6.....	133
Reconstruction of the Pyramid Dome geometry. a) dashed white line illustrates the former ideal crater wall position; the solid line marks the inferred dome outline; the black dot is the assumed vent location at the break in slope. b) view of the remnant summit dome from the W. c) top view of the combined paraboloid with a composite elliptical base; dimensions of elliptical radii are given. d) side view of the inferred dome geometry; the dome remnants are in dark grey; the inferred underlying slope on the upper flank is estimated to be c.20°.	
Figure 5-7.....	135
Hornblende types. a) type 1 with continuous reaction rim. b) type 2 with discontinuous reaction rim; present only in sample SD1; note individual Fe-Ti oxide crystals are visible. c) type 3 partial to fully replaced hornblende crystals. d) type 1 with observed brown glass fringing the reaction rim; only observed in sample SD6. Scale bar is 100 µm in a) otherwise 25 µm.	
Figure 5-8.....	136
Histogram of type 1 hornblende reaction rim thicknesses averaged for crystals and entire samples.	

Figure 5-9.....	137
Hornblende compositions of the summit lava dome, Mt. Taranaki. a) Na+K (A-site) vs. Al ^{IV} . b) Mg# vs. Si; (c.p.f.–cation per formula unit). For comparison are shown recalculated pargasitic hornblende compositions of Unzen volcano, Japan (Sato et al., 1999; Browne et al., 2006; Nakada and Motomura, 1999) [Unzen matrix refers to groundmass crystals], Soufrière Hills Volcano, Montserrat (Barclay et al., 1998; Rutherford and Devine, 2003), Mt. St. Helens, USA (Rutherford and Hill, 1993), Colima, Mexico (Luhr, 2002) and Cerro la Pilita, Mexico (Barclay and Carmichael, 2004).	
Figure 5-10.....	138
Compositions of Fe-Ti oxide phenocrysts and inclusions in clinopyroxene and hornblende. a) Al vs. Ti. b) Ti/Al vs. Fe ³⁺ # (c.p.f.–cations per formula unit). Cations are calculated on the basis of 32 oxygens.	
Figure 5-11.....	139
Glass compositions of inclusions in clinopyroxene and hornblende. Silica is used as differentiation index. Note that inclusions in different host minerals form separate groups.	
Figure 5-12.....	140
Bulk rock compositions of the summit lava dome and the Tahurangi BAF A and B deposits. Dome compositions are distinct to Tahurangi rocks as illustrated for Al ₂ O ₃ (a), Mg# (b, d) Fe ₂ O ₃ (e) and Zr (f).	
Figure 5-13.....	141
Trace element patterns of the Pyramid Dome and Tahurangi BAF deposits normalised to N-MORB (a) and chondrite (b). Pyramid Dome rocks and Tahurangi BAF deposits show nearly identical trace element patterns. For the light rare earth elements slightly higher abundances in Tahurangi BAF deposits are noted. Normalisation after Sun and McDonough (1989).	
Figure 5-14.....	142
Texture of rock sample SD6. a) photograph shows sub-vertical, near parallel crack patterns and cavities. b) modified image of a) highlighting cracks and cavities in black.	
Figure 5-15.....	144
Lava dome growth patterns are illustrated in a-e. Exogenous and endogenous dome growth occurred simultaneously. f) demonstration of the inferred exogenously (dark grey) and endogenously (light grey) formed surfaces as observed on the dome remnants.	
Figure 5-16.....	147
Reconstruction of the summit dome failure. a) erosion scars on upper flank (white arrows) as well as the curvatures of the amphitheatre, and the scars to the SW and S define the geometry of individual collapse sectors. b) dome geometry with individual sectors I-IV and their flow directions. c) cross-section of the dome showing the dome remnants (grey), and the disintegration of dome rocks along listric faults.	
Figure 5-17.....	149
Aluminium-in-hornblende geobarometer shown as histogram for hornblende phenocrysts (core and rim) and microphenocrysts. Calculated after Johnson and Rutherford (1989) and corrected by -1.5 kbar.	
Figure 5-18.....	150
Comparison of calculated hornblende crystallisation pressures for various Mt. Taranaki rocks and xenoliths. Granodiorite xenoliths contain all required mineral phases for the Al-in-hornblende geobarometer (Johnson and Rutherford, 1989) and therefore were not corrected. Note that some hornblende crystals of hornblende gabbros and hornblende-pyroxene gabbros indicate crystallisation below (<1 kbar) the inferred hornblende stability limit.	

Figure 6-1.....	179
<p>Mount Taranaki (lower right) has produced mainly lava dome eruptions in the past 800 years with Block-and-Ash Flow deposits making up the fan between the Maero and Pyramid Stream and in the Hangatahua River (BAF – Block-and-Ash Flow, ppf – pumice pyroclastic flow). Star (top left) indicates the most distal outcrop discussed in the text. To the NNW of Mt. Taranaki are the south flanks of Pouakai volcano. The inset shows the Taranaki peninsula with the Taranaki Volcanic Lineament (SLI-Sugar Loaf Islands, K-Kaitake, P-Pouakai, T-Taranaki). Major onshore and offshore faults: IF-Inglewood Fault, MF-Manaia Fault, NF-Norfolk Fault, OF-Oanui Fault. Contours are 300 m. Modified after Sherburn and White (2005).</p>	
Figure 6-2.....	181
<p>Variations in bulk vesicularity and connected porosity of single clasts for grey pumice (a), banded pumice (b), and black and brown pumice (c). Crosses represent the range in vesicularity per clast using minimum, mean and maximum values (see inset in a). Variations in bulk vesicularity refer to single cores cut in two (ϕ_{core}) and overall clast variations with multiple cores (ϕ_{clast}). Note different scale in b). See text for details.</p>	
Figure 6-3.....	183
<p>Field photographs of a) succession of three pumice pyroclastic flow deposits on the upper south flanks; sketch shows a general assembly of pumice types and grey dense lithics, b) grey pumice clasts of unit 3, c) eroded surface into unit 2 showing the scattered grey pumice [1] from airfall, brown pumice [2], banded pale grey to dark brown pumice [4], and the dense fractured andesite clasts [L]; d) lower contact of a distal BAF deposit, c.13.5 km from source (star in Fig. 1). See text for field description.</p>	
Figure 6-4.....	184
<p>Distribution of Burrell Lapilli deposits: a) isopachs in cm including the BAF deposit (black) to the NW for reference, black squares represent mapping locations for fall deposits only, P-Pouakai, contours 100 m; b) isopleths for pumice clasts and pumice pyroclastic flow deposits on the upper flanks (black), see inset in a) for location; c) isopleths for lithic clasts, same outline as in b). Numbers in b) and c) are clast diameters in cm.</p>	
Figure 6-5.....	187
<p>Thin-section photographs illustrating basic vesicularity differences of juvenile clasts. a) dense grey lithic, b) semi-vesicular black pumice, c) vesicular black pumice with isolated and coalesced vesicles, crosses mark plagioclase crystals; d) grey pumice with isolated large single vesicles as well as larger coalesced vesicle. Scale bar is 100 μm in a-c and 500 μm in d. See text for details.</p>	
Figure 6-6.....	188
<p>SEM images of grey (a) and brown (b) pumice (note a and b are binarised; black=vesicles, white=glass + crystals); c) shows a large coalesced vesicle; d) consists of three SEM images showing the transition from grey to brown in banded pumice. Scale bar is 10 μm in c), otherwise 100 μm.</p>	
Figure 6-7.....	189
<p>Hornblende reaction textures in different clast types: a) fresh hornblende with no reaction rims in pumice, b) single Fe-Ti oxide crystals are attached to the hornblende rims in black semi-vesicular pumice, c) hornblende in dense grey lithic clast shows resorption textures and is partially replaced by clinopyroxene, plagioclase and Fe-Ti oxide crystals or is fully replaced (lower left); note abundant plagioclase microlites in groundmass. Scale bar is 100 μm.</p>	
Figure 6-8.....	190
<p>Bulk rock geochemistry of pumice and grey andesite clasts in a multi element oxide vs. SiO_2 diagram. The calculated fractionation trend pumice – grey lithics is in good agreement for the majority of clasts (solid line) with some variation for the most evolved clast (dashed line). See Table 6-3 for details.</p>	

Fig. 6-9.....	194
Groundmass glass compositions of pumice types presented in the Al ₂ O ₃ vs. SiO ₂ diagram. Modelled glass composition changes due to plagioclase and clinopyroxene crystallisation (thick solid line) and is in good agreement with linear regression line (thin solid line). The small inset shows six data points of one lapillus (SD20) demonstrating relative glass homogeneity. The dimensions of the box are 1 wt. % for Al ₂ O ₃ and SiO ₂ .	
Figure 6-10.....	196
Connected porosity vs. bulk vesicularity of all pumice types. Brackets represent 95% confidence limit for the mean of each pumice population. Solid lines represent 0% and 10% and dashed line 5% isolated pore volume.	
Figure 6-11.....	197
Connected and bulk vesicularity vs. permeability. a) data of this study with upper and lower data limits (black lines) of $y=5\times 10^{-19} x^{-4.5314}$ and $y=6\times 10^{-21} x^{-4.5314}$, respectively. Note there are six specimens with three cores cut in three mutually perpendicular directions. Upper inset shows cores cut in two perpendicular directions. b) comparison of our data with published literature: Montserrat (Melnik and Sparks, 2002), Big and Little Glass Mountains (Rust and Cashman, 2004), Pichincha (Wright et al., 2007); grey lines are limits of Klug and Cashman (1996).	
Figure 6-12.....	202
Reconstruction of eruptive events during the Burrell Lapilli eruption. Changes in bulk rock silica contents are illustrated in Stage a. Bubble nucleation levels 1-3 correspond with erupted units 1-3. See text for further details.	
Figure 6-13.....	213
Correlation of pyroclastic flow deposits associated with the Burrell episode. Note that the major unit from medial to distal represents the Burrell Breccia (A). In section S04-133, the thin pyroclastic pumice flow deposits represent Burrell Breccia (B) units 1-3. Exposures are sorted by stream and planimetric distance from source. See figure legend for further details (RAD – rock-avalanche deposit). Magnetism was measured by a portable fluxgate magnetometer. For clarity, older exposed units were omitted. Outcrop numbers and profiles refer to Platz (2001) except S04-133. For list of samples and coordinates of outcrops see Appendix B.	
Figure 6-14.....	215
Estimates of total water and carbon dioxide contents in melt inclusions. a) total H ₂ O at 3550 cm ⁻¹ vs. molecular H ₂ O at 1630 cm ⁻¹ , and b) molecular CO ₂ at 2350 cm ⁻¹ vs. total H ₂ O at 3550 cm ⁻¹ .	
Figure 6-15.....	216
Different shapes of melt inclusions in clinopyroxene. a) overview of crystal 2 (sample SD20); note the many inclusions of glass, plagioclase, apatite and Fe-Ti oxide which are mostly oriented along crystallographic planes, b) two close-ups as marked in a); I - represents common but very small melt inclusions found in clinopyroxene, which are unsuitable for FTIR analysis. Their shape is near spherical to ovate; II - the bottle-neck shape is typical for leaked melt inclusions, c) overview of crystal 11 (sample SD32) showing a large irregular shaped melt inclusion, d) close-up of c) showing the impossibility of using these inclusions for FTIR analysis; it can be assumed that the inclusion extends further into the crystals as indicated by the diffuse outline of the melt inclusions further to the right, e) section of crystal 4 (sample SD20) showing two types of melt inclusions, the reddish-brown coloured inclusions are probably altered in comparison to the brown inclusions to the right; note again the irregular outline of the inclusions, f) section of crystal 8 (sample SD9D); abundant sheet-like inclusions probably oriented along crystallographic planes; the inclusions around the Fe-Ti oxide inclusions (black) appear to be connected. Scale bars are 100 μm in a), c), and e), 50 μm in d) and f), and 10 μm in b).	

Figure 6-16.....	218
Preliminary results of the thermal analysis studies. a) trial and error series of sample SD20, b) reproducibility results of different pumice samples; it is noted that for the same sample the maximum weight loss is often observed at similar temperatures.	
Figure 7-1.....	229
Groundmass texture (a-d) and crystallinity (e-h) of clasts from pyroclastic flow deposits. a and c) two groundmass glasses, b-d) differences in degree of groundmass crystallisation in clear translucent and brown glasses, e) semi- to hyaline, clear translucent glass; note microvesicularity, f) same image as in e) under crossed polarised light, g) semi- to holocrystalline brown glass, h) same image as in g) under crossed polarised light.	
Figure 7-2.....	231
Bulk rock composition of selected Maero eruptives. Block-and-Ash Flow deposits are not differentiated and the Pyramid Dome and the Turtle are omitted for clarity. For comparison, selected lava flows of the upper main cone and Fanthams Peak are plotted. $Mg\# = 100[Mg^{2+}/(Mg^{2+}+Fe^{2+})]$; all iron as Fe^{2+} .	
Figure 7-3.....	233
Analytical totals of all EMP glass analyses (a) and sample averages (b) are plotted against silica content. Estimated glass water contents using the water-by-difference method (WBD) are shown on the right axis. The terms andesite, dacite, and rhyolite refer to the TAS-classification scheme of Le Maitre et al. (1989).	
Figure 7-4.....	236
Calculated melt viscosities, η , are plotted against silica abundances. Values of the models of Shaw (1972) and Hui and Zhang (2007) are shown for H_2O contents of 0.1 wt.%, 1 wt.% and WBD at $T=900\text{ }^\circ\text{C}$ and $P=1$ bar. Solid and dashed lines are regression lines of η at WBD for the Hui-and-Zhang- and Shaw-models, respectively.	
Figure 7-5.....	237
Calculated magma viscosities, η_a , are plotted against SiO_2 contents. Lower and upper crystal volume fractions of 30% (a) and 55% (b), respectively are used for the calculation based on the calculated melt viscosities (see Fig. 7-4). Viscosities are calculated using H_2O contents of 0.1 wt.%, 1 wt.% and WBD at constant $T=900\text{ }^\circ\text{C}$ and $P=1$ bar. Solid and dashed lines are regression lines of η_a at WBD for the Hui-and-Zhang- and Shaw-models, respectively.	
Figure 7-6.....	238
Bulk SiO_2 contents (a) and $Mg\#$ (b) are plotted against K_2O in chronological appearance of eruption episodes.	
Figure 7-7.....	241
Calculated Mt. Taranaki melt viscosities are compared to calculated melt viscosities of Merapi volcano (Indonesia), Soufrière volcano (St. Vincent), and Soufrière Hills Volcano (Montserrat), using the same parameters. Solid lines are regression lines for Taranaki data.	

Figure 7-8.....242
Calculated Mt. Taranaki magma viscosities plotted against SiO₂ are compared to other andesite to rhyolite volcanoes. Since Taranaki viscosity calculations are based on glass chemical compositions of the Maero Eruptive Period, the range in bulk silica contents of rocks erupted during this period are used to allow comparison to published data. Upper and lower viscosity abundances are taken from Fig. 7-5. Taranaki data are illustrated by two parallelograms with upper and lower limits representing crystal volume fractions of 55% and 30%, respectively. The grey parallelogram corresponds to 1 wt.% melt water content, whereas the dashed parallelogram relates to water contents determined by WBD. Data source: silicic lava flows (Murase and McBirney, 1973; Fink, 1980; Navarro-Ochoa et al., 2002; Manley, 1996; Harris et al., 2004; and McKay et al., 1998); Mt. St. Helens (Murase et al., 1985; Scandone and Malone, 1985); Unzen volcano (Suto et al., 1993; Goto, 1999; Sato et al., 1999); Soufrière Hills Volcano, Montserrat (Voight et al., 1999; Sparks et al., 2000); Soufrière volcano, St. Vincent (Huppert et al., 1982); Merapi volcano (Siswamidjono et al., 1995).